## The paraglacial, giant landslide dam of Phoksumdo Lake (Dolpo District, Western Nepal Himalaya).

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Large catastrophic slope failures have recently retained much attention in the northern dry Himalayas and Central Asia [1]. Their occurrence is generally interpreted as resulting from three main potential triggers: earthquakes, post-glacial debuttressing, permafrost melting. The Phoksumdo lake (3600 m asl) is the second largest lake in Nepal (area of 4.5-to-5 km<sup>2</sup>). It owes its origin to the damming of the Suli Gad River by a giant (4.5 km<sup>3</sup>) collapse of a mountain wall culminating at 5148 m asl, SE of the lake [2]. The landslide morphology presents a series of mounds and depressions of varying size (a few-to-ten metres high or deep). The landslide mass is composed of large blocks of massive, palaeozoic dolomites found as far as west of Ringmo Village and down to the southern rim of the landslide dam [3] [Fig. 1]. The coarse debris is widely mantled by a thick (up to 10 m) cover of silts, well prominent downstream above the Suli Gad gorges. Downstream, most of the Suli scarp is upheld by limestone's outcrops, but along the central part, where the former Phoksumdo gorge is filled in by dark grey, fine sediments (till material) [2,3], and overlain by orange conglomerates (including dolomites boulders), then by vellowish, unsorted conglomeratic material composed of blocks and calcareous silty matrix [3]. This indicates the damming of the lake was caused by rockslides bodies derived from different parts of the rocky face bounding the lake in its SE part. On the basis of reddish buried soils in silts, Yagi [2] considered rockslide occurrence at ≈ 30-40 ka following glacial retreat.



Figure 1. Left: The giant collapse (4.5 km<sup>3</sup>) of the mountain wall (5148 m asl) made of Dhaulagiri limestones is at the origin of the Phoksumdo lake. Right: View of the rockslide mass and location of dating samples.

In order to test this assumption, we carried out AMS measurements of the  $^{36}$ Cl concentration of the samples from the main scarp [Fig. 1]. Measurements were performed at the 5 MV national AMS facility ASTER located at CEREGE, Aix en Provence, France. All  $^{36}$ Cl concentrations were normalized to the KNSTD1600 calibration material ( $^{36}$ Cl/ $^{35}$ Cl =  $1.6*10^{-12}$  provided by K. Nishiizumi). The used decay constant of  $2.303 \pm 0.016*10^{-6}$  a<sup>-1</sup> corresponds to a  $^{36}$ Cl half-life ( $T_{1/2}$ ) of  $3.014*10^{5}$  years. Analytical uncertainties include the counting statistics, machine stability ( $\sim 0.5\%$ ) and blank correction ( $^{36}$ Cl/ $^{35}$ Cl blank ratio was on the order of  $\sim 10^{-14}$ ). Cosmic ray exposure ages derived from these  $^{36}$ Cl concentrations were calculated using production rates corrected for topographic shielding following [4] and scaled to the sites position (latitude, altitude) using the Stone scheme [5] from sea

level and high latitude production rates integrating all production pathways (spallation - including a spallation production rate of  $42.0 \pm 2.0$  atoms of  $^{36}\text{Cl g}^{-1}\text{Ca a}^{-1}$  at sea level and high latitude, thermal neutron and muon captures, radiogenic production) and the rock's chemical composition.

Sample ID	Sample weight	% CaO	<sup>36</sup> C1/ <sup>38</sup> C1 (10 <sup>34</sup> )	"CL/"CI	C1	<sup>36</sup> Cl	Exposure Age of surface (uncroded surface)	Uncertaintles
PHOK COS 1	67,29	55,18%	133,8	35,64	2,79	6,8	6987	566
PHOK COS 2	69,08	54,70%	382,0	24,08	4,58	19,8	20882	1687
PHOK COS 3	72,015	54,59%	314.8	10,42	13,82	19,5	20888	1661

Table 1: AMS measurements of the <sup>36</sup>Cl concentration of the samples from the main scarp.

The post glacial origin is now confirmed by cosmogenic dating: two consistent  $^{36}$ Cl ages of 20,885 ±1675 [Tabl.1] argue for a single, massive event of paraglacial origin, that fit well with the last chronologies available on the Last Glacial Maximum in the Nepal Himalaya [6].

The IRSL samples (PHOK 2 and PHOK 1) were collected respectively at the top and the bottom of the silty deposit, with the reddish soil between. IRSL measurements were performed on alkali feldspar grains (60-80 mm). These were extracted by dry sieving, then treated with HCl and  $H_2O_2$  to remove calcium carbonate and organic matter. The heavy liquid isolation of alkali feldspars (density <2.58 s.g.) was achieved by using sodium polytungstate. Palaeodose (De) determinations on alkali feldspars were done using the single-aliquot regenerative-b dose (SAR) method, following the Montreal laboratory protocol on very small aliquots (about 50 grains per aliquot). This protocol [7] has been applied to correct for the observed fading. The corrected IRSL ages of the samples, respectively  $4.7 \pm 0.3$  ka and  $12.6 \pm 0.7$  ka, are stratigraphically consistent. They suggest a progressive trend to dryness of these Northern Himalayan regions. The high lime content of the landslide components explains the dam stability and the persistence of the sacred lake of Phoksumdo.

As many other examples this mega-collapse confirms the fact that giant landslides play a significant role in both the destruction of Himalayan topography and in bedrock protection from river incision, hence delaying sediment transport outward from the mountain zone [1, 8, 9, 10].

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